SPECTRAL AND PHOTOMETRIC CHARACTERISTICS OF MID-LATITUDE AURORA DURING THE MAGNETIC STORM OF MARCH 17, 2015

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Abstract. We study the spatiotemporal dynamics of mid-latitude aurora from observations in the south of Eastern Siberia during St. Patrick’s severe geomagnetic storm on March 17, 2015. We perform a morphological analysis of characteristics of the observed auroras. A preliminary conclusion is drawn that the analyzed event is the result of the manifestation of mid-latitude auroras of two types (type “d” and SAR arc) and ordinary aurora observed at the northern horizon. The maximum intensity of the dominant emission [OI] at 630.0 nm (~14 kR) allows this mid-latitude aurora to be attributed to the extreme auroras occurring in mid-latitudes, which is second only to the November 20, 2003 superstorm (~19 kR).

Keywords: geomagnetic storm, mid-latitude aurora, spectrum of mid-latitude aurora, 557.7 and 630.0 nm emissions.

INTRODUCTION

Mid-latitude auroras (MLA) are a relatively rare geophysical phenomenon. According to [Krakovetsky et al., 1989], the probability of MLA observation in the Northern Hemisphere is five cases per year at geographical latitudes 55°–60° N; one, at 50°–55° N; 0.1, at 40°–50° N. The main dominant emission of MLA is the forbidden 630.0 nm atomic oxygen [OI] line whose intensity depends on the degree of development of the ring current (Dst index) during a magnetic storm (MS) [Troutt, 1973; Rassoul et al., 1992; Mikhailov et al., 2004]. The minimum Dst index (Dstmin) for the March 17, 2015 MS is ~222 nT [http://wdc.kugi.kyoto-u.ac.jp/dst_realtime]. This, according to [Gonzalez et al., 1999], allows us to assign this storm to extreme MS. The last MS with Dst ≤ ~220 nT occurred on May 15, 2005 in solar cycle 23. Thus, the March 17, 2015 MS is one of the strongest (by Dst) magnetic storms of solar cycle 24.

This paper reports the results of MLA observations made with a complex of optical instruments in the south of Eastern Siberia during St. Patrick’s geomagnetic storm on March 17, 2015. We perform a preliminary morphological analysis of characteristics of the observed MLA, without detailed, with some exceptions, comparison with isospheric, magnetospheric, satellite data and models. Such a comparison that is undoubtedly of great interest may be a subject for further research.

EQUIPMENT AND OBSERVATIONAL METHOD

The observations have been made at the ISTP SB RAS Geophysical Observatory (GO) (52° N, 103° E), located in the Tunka valley (Tory) in the south of Eastern Siberia. During the observation of the March 17, 2015 magnetic storm, the complex of optical instruments included two spectrographs, a wide-angle color camera, an all-sky camera at 630 nm, and a Fabry—Perot interferometer (630 nm). All the optical instruments operated in automatic mode.

Spectrographs. A patrol spectrograph with low spectral resolution is designed to detect viewing spectra of nightglow [http://atmos.iszf.irk.ru/ru/data/spectr]. During the March 17, 2015 MS, the spectrograph was oriented northward with a ~23° elevation of the optical axis above the horizon; the exposure time was 260 s. The absolute calibration of the patrol spectrograph is described by Mikhailov [2018]. The Shamrock SR-303I spectrograph [Beletsky et al., 2016] was also oriented northward with a ~30° elevation of the optical axis above the horizon. Measurements have been made in series of 30 frames with 1 s exposure time (frames of each series are summed up). The full operating spectral band of the Shamrock SR-303I spectrograph is 492–896 nm.

The wide-angle color camera is designed to detect and examine spatio-temporal variations of nightglow and to monitor atmospheric transparency [http://atmos.iszf.irk.ru/ru/data/color]. The camera was focused on the northern sky, on the celestial pole. Its angular field of view is ~85°. The exposure time is 300 s.

The all-sky camera KEO Sentinel is designed to record spatial patterns of 630 nm emission intensity (an emission height is 180–300 km). The half-width of the interference filter is ~2 nm. The viewing direction is zenith, the field of view is 145°, the exposure time is 30 s.

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Keto Scientific’s Fabry—Perot interferometer. Characteristics of the Fabry—Perot interferometer (FPI) in hand are discussed by Shiokawa et al. [2012]. The analysis of interferograms to obtain variations in line intensities is based on the method set forth by Harding et al. [2014]; and Vasilyev et al. [2017] give details of adaptation of this method for a particular device and observing program in use. During the observation of the magnetic storm, FPI worked in the mode of patrol survey of five directions – cardinal directions and zenith.

OBSERVATIONAL RESULTS AND DISCUSSION

The March 17, 2015 geomagnetic storm, named according to the day of its onset St. Patrick’s storm, is one of the severest geomagnetic storms of solar cycle 24; and by the Dst index, the most powerful storm in this cycle. The geomagnetic effect of this storm is due to the superposition of two consecutive storms [Kamide, Kusano, 2015].

Features of the spectral composition of the March 17, 2015 MS It is believed that in MLA spectra, lines and bands can appear or intensify which are typical of both mid-latitude airglow and auroras, in particular, 557.7 and 630.0 nm atomic oxygen [OI] lines, 520 nm nitrogen [NI] line, bands of the first negative system \( N_2^+ \) 391.4, 427.8, and 470.9 nm, etc. In terms of intensities and ratio of dominant emissions, MLA falls into different types [Rassoul et al., 1993].

Figure 1 shows a typical spectrum in the 500–700 nm wavelength range, recorded by the Shamrock SR-303I spectrograph during MS.

Figure 2 presents the dynamics of the 630.0 and 557.7 nm dominant emissions of March 17, 2015 and their ratio, derived from patrol spectrograph observations. For comparison, the Figure depicts the behavior of the Dst index over the same period.

According to GO observations of MLA during severe MS in previous solar cycles [Mikhalev et al., 2004, 2005], it was concluded that depending on the level of geomagnetic activity and magnetic storm phase MLA can be assigned to different types: «d» (with the 630.0 nm dominant emission), SAR arc (630.0 nm), and aurora (557.7 and 630.0 nm). MLA of these types differ in spectral composition, intensities of dominant emissions, and detection probability at night, thus reflecting the current state and dynamics of magnetospheric-ionospheric structures and their projections on heights of the upper atmosphere.

As derived from GO optical observations over the period 1991–2017, by maximum intensity of the 630.0 nm dominant emission (~14 kR) the MLA of interest is second only to the MLA occurring during the November 20, 2003 superstorm with \( Dst_{min} = -465 \) nT (~19 kR). During other strong MS with a lower Dst index, intensities of the 630.0 nm dominant emission do not exceed 10.5 kR (e.g., during the March 24, 1991 MS, \( Dst_{min} = -281 \); during the April 6, 2000 MS, \( Dst_{min} = -287 \); during October 30, 2003 MS, \( Dst_{min} = -401 \) [Mikhalev et al., 2005]). Other features atypical of the MLA previously observed at GO are the lack of dependence of the 630.0 nm emission on Dst at night and the presence of pronounced bursts of the 630.0 nm emission with characteristic durations of ~1 hr, more typical of substorm activity in subauroral and auroral latitudes. In the emission spectrum of the analyzed MLA, in addition to the 557.7 and 630.0 nm dominant auroral emissions also typical of the mid-latitude atmospheric emission, we recorded the 520 nm [NI] emission and the \( N_2^+ \) 586.7 nm band. Thus, by its spectral characteristics the MLA observed in the northern sky may be assigned to the type «d» with manifestations of substorm activity superimposed in separate time intervals.

Figure 1. MLA spectrum in the 500–700 nm wavelength range, recorded by the Shamrock SR-303I spectrograph during the March 17, 2015 MS

Figure 2. Intensities of the 630.0 and 557.7 nm emissions, and Dst variations during the March 17, 2015 magnetic storm
Photometric aspect observations made with the color wide-angle camera in RGB channels. As noted above, the color camera during the March 17, 2015 MS was pointed at the north. Figure 3 presents images of the northern sky in R, G, and B channels, captured by the camera at 15:22 UT.

The diffuse brightness distributions (Figure 3) occupying most of the northern sky in the R channel and localized near the horizon in the G channel lasted almost throughout the night on March 17, 2015. Using spectral characteristics of R, G, and B channels and features of the spectral composition of the upper atmosphere emission during MLA (Figures 1, 2), within certain assumptions we can estimate the southern boundary of precipitation during MS under study. Thus, Podlesny, Mikhailov [2015] and Mikhailov et al. [2016] have shown that disturbances in the R channel of the color camera during the March 17, 2015 MS are largely caused by increased intensity of the 630.0 nm emission. Suppose that the disturbance in the G channel is similarly related to the enhanced 557.7 nm emission, then spatio-temporal variations in brightnesses of R and G images may be caused by variations in the 630.0 nm and 557.7 nm emission intensities respectively. It is believed that the 557.7 nm emission with characteristic heights ~90–110 km is dominant in aurora [Omholt, 1974]. Thus, knowing the angular height of the maximum emission near the horizon in the G channel, which in our case is ~8°–10°, we can estimate the distance to the emission region. Referring to the obtained data, the distance to the visible precipitation zone can be estimated as ~530–785 km. This corresponds to the displacement northward of the observation point by ~5°–7° from the ISTP SB RAS GO geomagnetic latitude (47.9° N).

MLA images captured by the color camera in the R channel have much larger angular sizes than those taken in the G channel. Besides the region coinciding with airglow in the G channel at angular heights 8°–10°, there is an intense airglow in the angular range ~10°–35° above the horizon with a peak intensity of ~22°.

As already mentioned, the disturbances in the R channel of the color camera during the March 17, 2015 MS may be associated with increased intensity of the 630.0 nm emission [Podlesny, Mikhailov 2015; Mikhailov et al., 2016]. In this case, assuming that the 630.0 nm emission region coincides with the 557.7 nm emission region with an estimated range ~530–785 km from the observation point, we can assess the vertical distribution of the 630.0 nm emission during MLA. Thus, the maximum 630.0 nm emission is at heights ~200–300 km, and the upper boundary is at 350–550 km. The large height of the maximum 630.0 nm emission and the vertical extent of MLA in the 630.0 nm emission, as largely the diffuse nature of airglow, allows us to assign this aurora by color characteristics to red auroras of type "d" (red color of the total aurora is typical of subauroral auroras, which are characterized by large (hundreds of kilometers) heights). At the same time, in some short intervals of observations (e.g., at 16:30–18:30 UT with increasing 557.7 nm emission), this aurora can be qualified as a type "a" aurora (red color in the upper part of the aurora and prevailing green color in the lower part). It is believed that a direct source of the 630.0 nm emission (the "D level") in type "d" MLA is electrons with energies ~10–1000 eV [Rasoul et al., 1993]. Thus, data from the color camera can supplement spectral observations, for example, suggest that by its color characteristics the March 17, 2015 MLA can be attributed to the red type "d" aurora and, in some intervals, to the type "a" aurora; they can also enable us to estimate the distance to the equatorial boundary of the detected MLA, which is ~530–785 km, being equivalent to geomagnetic latitudes ~47°–49°.

Photometric observations in the 630.0 nm emission with the all-sky camera KEO. The angular field of view of the all-sky camera KEO is ~145°. When pointed at the zenith, it captures an image at angular heights >18.5° above the horizon, thus allowing us to compare the dynamics of brightnesses of images in a northerly direction with spectrograph data. In Figure 4 are images in the 630 nm emission captured by this camera at 12:53, 15:30, 16:39, 19:00, and 21:08 UT versus the dynamics of brightnesses of R channel obtained by the patrol spectrograph.

The analysis of the images in the 630.0 nm emission (Figure 4) allows us to identify several specific development stages in the analyzed MLA. The first stage (12:15–13:00 UT) features afterdusk undisturbed conditions.

In Figure 4, to this stage corresponds the first left image (the near-horizon brightness enhancement in the form of a circle is due to the Van Rine effect, caused by an increase in the optical thickness of the luminescent layer toward the horizon). The second stage (~13:00–13:30 UT) is associated with the appearance of a diffuse spot near the northern horizon, which moves along the horizon. During the next stage (~13:45–15:30 UT), the second layer of weak airglow is formed (equatorward of the diffuse spot) which moves slowly to the equator. From 15:30 UT and to the end of observations near the northern horizon.

Figure 3. Images of the northern sky in the R, G, and B channels, captured by the color camera at ISTP SB RAS GO at 15:22:30 UT (a); keogram of MLA in the R channel (b)
horizon, airglow spots varying in shape are continuously recorded (an image second from the left in Figure 4), which are probably a manifestation of ordinary auroras.

The second layer of airglow, varying in shape (sometimes with a wavy equatorial boundary) and intensity, in the near-zenith region takes on a form of arc resembling the known SAR arcs [Rees, Roble, 1975], which, slowly moving equatorward, becomes more pronounced (Figure 4). In some periods, we can identify the appearance of the second, less weak arc, located to the north of the main arc (Figure 4, the first right image).

The detected arc moves to the south with an average velocity of ~20 m/s (calculated for an emission height of ~400 km). The movement of the arc in latitude correlates well with Dst variations. The velocity of displacement and size of this arc match the parameters of SAR arcs reported in [Ievenko, Alekseev, 2004].

The interpretation of the images, presented in Figure 4, as a SAR-arc phenomenon is supported by the fact that during this magnetic storm the subauroral optical station Maimaga in Yakutia (63° N, 129.5° E) in the evening MLT sector also recorded an intense SAR arc [Ievenko, Parnikov, 2017].

The keogram for ~15:00–15:30 UT (Figure 3) also exhibits a structure moving from smaller to larger angular heights and corresponding to times of beginning of the formation of the SAR arc as derived from data acquired with the all-sky camera (see, e.g., the image second from the left in Figure 4).

According to the classification [Rassoul et al., 1993], SAR arcs are mid-latitude auroras. It is now established that SAR arcs along magnetic field lines are adjacent to the region of interaction between hot ring current particles and cold external plasmasphere and, most likely, “represent the radial gradient of cold plasma density, where plasmospheric electrons are heated by energetic ring current particles” [Ievenko, Alekseev, 2004]. It should be kept in mind that SAR arcs are often associated with the projection of the plasmapause and respectively with the equatorial wall of the main ionospheric trough (MIT) [Khorosheva, 1987]. It is believed that a direct source of the 630.0 nm emission (the D level) are electrons with energies of <10 eV [Rassoul et al., 1993].

Thus, the analysis of observations made with the all-sky camera in the 630.0 nm emission shows that throughout the night near the northern horizon there were airglow spots with continuously varying shape, which are likely to be a manifestation of ordinary auroras.

Within the interval ~16–21 UT in the near-zenith region, there was a quite intense SAR arc. This may indicate the displacement of the MIT equatorial wall to the geomagnetic latitude of ~47.9° N in the longitudinal sector of interest during this storm.

**Interferometric observations in the 630.0 nm emission**

Figure 5 presents the results of measurements of the 630.0 nm emission intensity and Doppler temperature T_{630} in three directions—north, zenith, south—at zenith angles of 60°, 0°, and 60° respectively. The northerly viewing direction of FPI and spectrographs is virtually the same, hence the dynamics of the 630.0 nm emission intensity in this direction is also nearly identical (Figure 2). It should be noted (Figure 5) that there is a large intensity gradient simultaneously between the northerly direction and the direction to the zenith and south and almost the same average Doppler temperatures T_{630} ~1450–1650 K after 15 UT in these three directions. Heating of the upper atmosphere during geomagnetic storms and an increase in the Doppler temperature measured from the 630.0 nm emission are well known [Truttse, 1973; Earle et al., 2013]. Among possible causes of the increase in the temperature of the neutral atmosphere during geomagnetic storms are currently considered to be the Joule dissipation of currents in the auroral zone at heights of the lower thermosphere (~100–140 km), followed by equatorward transport of the disturbed region due to additional circulation [Danilov, 2013; Earle et al., 2013]. In the subauroral ionosphere at heights above 200 km there may be frictional heating in a wide band of subauroral ion drifts westward near the plasmapause projection at F-layer heights. The drifts are most pronounced during substorms against the background of large-scale plasma convection [Zhang, 2017]. FPI data showing an increase in T_{630} for all viewing directions from ~1100–1150 K at the beginning of the measurements to ~1450–1650 K at 15–16 UT agree.
with experimental data and existing ideas about possible mechanisms of the neutral atmosphere heating at the F-layer heights during geomagnetic storms. However, the set of the obtained optical data, including temperature data, requires additional, more detailed studies of some features of the March 17, 2015 magnetic storm.

CONCLUSIONS

Nowadays there is no generally accepted concept of mid-latitude auroras in the literature. Some authors use the term low-latitude aurora, referring to auroral emissions observed during MS at geomagnetic latitudes of \( \leq 50^\circ \). At the same time, other authors adopt the term mid-latitude aurora, using as distinctive features the presence of \( \text{N}_2^+ \) in the emission spectrum and the high ratio \( r \) of the 630.0 and 557.7 nm emission intensities [Rassoul et al., 1993]. According to [Rassoul et al., 1993], MLA have some distinguishing features that allow us to separate them from ordinary auroras. Thus, according to the classification [Rassoul et al., 1993], MLA caused by electron precipitation are characterized by the ratio of the 630.0 nm red line intensity to the 557.7 nm green line intensity \( r > 10 \) for SAR arcs and \( 1 < r < 10 \) for type «d» MLA, whereas for ordinary auroras this value is usually within \( -0.2 \)–\( 1 \) [Omholt, 1974]. Spectral measurements made with spectrographs at narrow angles above the horizon during the March 17, 2015 aurora of interest do not allow us to correctly determine \( r \) because of possible latitudinal spacing of 630.0 and 557.7 nm emission layers. Nevertheless, the high 630.0 nm emission intensity, which is much higher than the 557.7 nm emission intensity, and the dominance of signals in the red channel over signals in the green channel as derived from data acquired with the color wide-angle camera at large angular heights indirectly indicate a rather high value of \( r(>1) \). This may be an argument for identifying this aurora as mid-latitude.

Another feature of MLA noted in many works [Truttse, 1973; Rassoul et al., 1992; Mikhail, 2001; Mikhail et al., 2004] is the pronounced dependence of the main dominant emission of MLA at 630.0 nm on the degree of development of ring current during MS (by \( Dst \)). This is associated with some differences in formation mechanisms between mid-latitude aurora and aurora polaris. In our case, this dependence is broken by intense bursts of the 630 nm emission, which likely have substorm nature. This suggests the possibility of interpreting this aurora also as ordinary aurora polaris. Having regard to the above, we can assume that during the main phase of the March 17, 2015 GO simultaneously recorded MLA and ordinary auroras; therefore, the term mid-latitude aurora in the title of this article is rather conventional.

Thus, the results of the study of MLA, observed in the south of Eastern Siberia during St. Patrick’s geomagnetic storm on March 17, 2015, allow us to draw a preliminary conclusion that this MLA is atypical of the longitude-latitude zone of interest. During the MS main phase, geophysical conditions might have been realized which allowed simultaneous recording of mid-latitude (type «d») and SAR arc and auroral (type “a”) forms of auroras.

Among the most interesting morphological features of the aurora, indicating, in particular, its atypical behavior, we can note the following:

- The maximum intensity of the dominant 630.0 nm emission (~14 kR) during the March 17, 2015 magnetic storm enables us to assign these MLA to the extreme MLA observed in the vicinity of GO in 1991–2017. By the manifestation in the 630.0 nm emission the analyzed MLA is second only to the period of the superstorm on November 20, 2003 (~19 kR) and is superior to auroras during MS with stronger geomagnetic disturbances (by \( Dst \)).
- There is no dependence of the 630.0 nm emission on \( Dst \) at night, which is atypical of MLA previously observed at GO.
- There are intense bursts of the 630.0 nm emission with typical durations of ~1 hr, which is more typical of substorm activity in subauroral and auroral latitudes.
- In the MLA spectrum, except for the dominant auroral 557.7 and 630.0 nm emissions also typical of mid-latitude airflow, we detected the 520 nm [NI] emission and the 586.7 nm \( \text{N}_2^+ \).
- Referring to interferometric observations in the 630.0 nm emission for ~3 hr there was an increase in the Doppler atomic oxygen temperature from ~1100–1150 to ~1450–1650 K.

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